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Non-linear wave equations for free surface flow over a bump

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22 **Abstract**

23 This study aims to develop a new wave equation model by modifying the Fully-nonlinear and strongly-
24 Dispersive Surface wave (FDS) equations. The modification was performed by applying a new expansion in a
25 series of the vertical coordinate, z'' , to the velocity potential while a simple expansion in a series of z was
26 applied to the FDS equations. Verification of the model was conducted by comparing with the theoretical
27 solutions of surface solitary waves. We applied the modified FDS equations to wave fields over a bump under
28 conditions with and without currents, which agreed very well with the time series of wave heights and velocity
29 obtained from laboratory experiments. The dispersion relationship computed using the normalized modified
30 FDS equations also agreed very well with the theoretical solution when we gave the number of expansion
31 terms as 3 with $\mu=2.5$. Additionally, the profile of surface waves computed with the modified FDS equation
32 was shown to have a larger width ridge, a bulbous-type wave, by comparing with a Trochoidal wave under the
33 condition of waves against a current.

34 *Keywords:* strongly dispersive, dispersion relationship, bulbous wave, laboratory experiment,
35 variational principle.

36

37 1. Introduction

38 In previous studies, Dean (1965) showed a significant effect of mean currents on the shape of surface waves.
39 Dalrymple (1974) showed the possibility for the existence of a surface wave with a larger width around a ridge
40 under a linear shear condition, which may be a similar condition to a constant current condition. Furthermore,
41 Teles and Peregrine (1988) demonstrated that such a large width ridge wave exists as a “bulbous” wave. The
42 application of a three-dimensional non-hydrostatic model is considered to be one of the solutions to analyze
43 surface waves with current effects playing a significant role in the deformation of surface waves. However,
44 since full three-dimensional non-hydrostatic model costs are very expensive compared to a layer model, it
45 seems necessary to develop a fully-nonlinear layer model that can include current effects but with cheaper
46 runtime.

47 The system of the Boussinesq-type equations is one of the methods often applied to solve such problems
48 (Mei and Mehaute, 1966; Perigrine, 1967; Madsen and Mei, 1969). An improved Boussinesq-type equation
49 can include a strong dispersive effect with nonlinearity and show high applicability to real scale phenomena
50 (Madsen et al., 1991; Madsen and Sorensen, 1992; Nwogu, 1993; Wei and Kirby, 1995; Wei et al., 1995;
51 Agnon et al., 1999). For example, FUNWAVE is a well-known model which can enable analysis of strong
52 nonlinearity in the near shore region (Wei and Kirby, 1995; Wei et al., 1995). Wei et al. (1999) developed a
53 method for generating waves in the Boussinesq-type equation, such as desired regular and random incident
54 wave with good agreements with laboratory experiments. As a different model application, Zaman et al. (2000)
55 applied the improved Boussinesq equation to analyze surface waves in Nakada Harbor of Okinawa Prefecture,
56 which demonstrated very good agreement with field observations.

57 However, strong nonlinearity may be important for the analysis of the effects of currents on waves.
58 Nakayama and Kakinuma (2010) developed the Fully-nonlinear and strongly-Dispersive Internal wave (FDI)
59 equation by following Isobe (1995) using the variational principle, which showed that the vertical profile of
60 horizontal velocity under a strong-nonlinear solitary wave can be evaluated accurately. Nakayama and
61 Kakinuma (2010) demonstrated that the FDI equation can reproduce a large amplitude solitary wave with
62 validation using the relationship between wave height and effective wavelength of a solitary wave (Koop and
63 Butler, 1981; Nakayama, 2006). Since the FDI equation was developed under multi-layer conditions
64 (Nakayama and Lamb, 2019), soliton resonance for internal solitary waves (Maxworthy, 1980; Miles, 1997ab;

65 Tanaka, 1993; Li et al., 2011) and the generation of solitary waves in a tilting tank (Horn et al., 2000; Horn et
66 al., 2001; Horn et al., 2002) can be analyzed successfully (Nakayama et al., 2019a). Furthermore, the solutions
67 obtained using the FDI equation have been applied to give boundary conditions for wave generation, which
68 has been verified from comparisons with the third-order solution of a solitary wave (Nakayama et al., 2012;
69 Nakayama et al., 2019b). Therefore, the Fully-nonlinear and strongly Dispersive Surface wave equations (FDS
70 equation) were developed based on the FDI equation using the characteristic of applicability to a multi-layer
71 system (Nakayama and Kakinuma, 2010; Nakayama et al. (2019a). In the FDI equation, velocity potential is
72 expanded in a series of z , which is used in the FDS equation. To improve the FDS equation, we made an
73 attempt to apply a new expansion in a series of z^μ . The improved model is called the modified FDS equation.

74 Firstly, we applied the modified FDS equation into solitary waves and made comparisons with theoretical
75 solutions by Grimshaw (1971) and Fenton (1972). Secondly, Thuy (2013) carried out laboratory experiments
76 in order to investigate the balance between waves and currents as morphology drivers. We thus applied the
77 modified FDS equation into their laboratory results under wave and wave-current conditions over a bump.
78 Two wave condition cases and one wave-current condition were compared with the modified FDS equation
79 and were verified using a timeseries of surface wave height and a time series of velocity over the bump. The
80 effect of μ was investigated by making comparisons between the FDS equation ($\mu = 1$) and the modified FDS
81 equation ($\mu > 1$) using the normalized modified FDS equation. The effect of an expansion number N on the
82 dispersion relationship was also investigated. Lastly, we investigated the profile of surface waves using Thuy's
83 wave-current laboratory experiments as a fundamental analysis, whether a bulbous wave exists or not under a
84 wave-current condition.

85

86 **2. Materials and Methods**

87 **2.1. FDS equation**

88 For the FDS equation (Nakayama et al., 2010; Nakayama et al., 2019a), velocity potential was used to
89 analyze waves using the variational principle of Luke (1967).

$$\mathbf{u} = \nabla\phi \text{ and } w = \frac{\partial\phi}{\partial z} \quad (1)$$

90 where \mathbf{u} is the velocity in the horizontal direction, ϕ is the velocity potential, w is the vertical velocity, and ∇
 91 is a partial differential operator in the horizontal plane, i.e., $\nabla = (\partial / \partial x, \partial / \partial z)$.

92 Incompressible fluids are assumed to be stable in still water (Fig. 1). The thickness in still water is denoted
 93 by h . The water surface is indicated as η . By following Isobe (1995) based on Luke (1967), the function for
 94 the variational problem, F , is determined by

$$F[\phi, \eta] = \int_{t_0}^{t_1} \int_A \int_b^\eta \left\{ \frac{\partial \phi}{\partial t} + \frac{1}{2} (\nabla \phi)^2 + \frac{1}{2} \left(\frac{\partial \phi}{\partial z} \right)^2 + gz + \frac{P}{\rho} \right\} \nabla \phi \quad (2)$$

$$P = -\rho gh \quad (3)$$

95 where g is gravitational acceleration; the plane A , which is the orthogonal projection of the object domain onto
 96 the $x - z$ plane is assumed to be independent of time.

97 In order to derive a set of equations, vertical integration is performed by expanding ϕ into a series in
 98 terms of α given a set of vertically distributed functions, Z_α , multiplied by their weightings, f_α .

$$\phi(x, y, t) = \sum_{\alpha=0}^{N-1} Z_\alpha f_\alpha(x, y, t) \quad (4)$$

$$Z_\alpha = z^\alpha \quad (5)$$

99 where N is the total number of an expanded function.

100 We substituted Eqs. (4)-(5) into Eq. (2), after the function was integrated vertically. Then the variational
 101 principle was applied to obtain the following Euler-Lagrange equations (Eqs. (6) and (7)).

$$Z_\alpha^\eta \frac{\partial \eta}{\partial t} - Z_\alpha^b \frac{\partial b}{\partial t} + \nabla \left(\int_b^\eta Z_\alpha Z_\beta dz \right) \nabla f_\beta - \int_b^\eta \frac{\partial Z_\alpha}{\partial z} \frac{\partial Z_\beta}{\partial z} dz f_\beta = 0 \quad (6)$$

$$Z_\beta^\eta \frac{\partial f_\beta}{\partial t} + \frac{1}{2} Z_\beta^\eta Z_\gamma^\eta \nabla f_\beta \nabla f_\gamma + \frac{1}{2} \frac{\partial Z_\beta^\eta}{\partial z} \frac{\partial Z_\gamma^\eta}{\partial z} f_\beta f_\gamma + g\eta + \frac{P}{\rho} = 0 \quad (7)$$

102 Finally, the FDS equations are

$$\eta^\alpha \frac{\partial \eta}{\partial t} + \frac{1}{\alpha + \beta + 1} \nabla (\eta^{\alpha+\beta+1} - b^{\alpha+\beta+1}) \nabla f_\beta - \frac{\alpha\beta}{\alpha + \beta - 1} \nabla (\eta^{\alpha+\beta-1} - b^{\alpha+\beta-1}) f_\beta = 0 \quad (8)$$

$$\eta^\beta \frac{\partial f_\beta}{\partial t} + \frac{1}{2} \eta^{\beta+\gamma} \nabla f_\beta \nabla f_\gamma + \frac{1}{2} \beta\gamma \eta^{\beta+\gamma-2} f_\beta f_\gamma + g\eta + \frac{P}{\rho} = 0 \quad (9)$$

103 where $\alpha = 0, 1, 2, \dots, N-1$, $\beta = 0, 1, 2, \dots, N-1$, and $\gamma = 0, 1, 2, \dots, N-1$.

104

105 **2.2. Modified FDS equation**

106 In the FDS equation, velocity potential is expanded in a series of z (Eq. (5)), which shows the vertical
 107 velocity as Eq. (10). Therefore, the flat bottom boundary condition, $w = 0$ at $z = 0$, cannot be satisfied.

$$w = \frac{\partial \phi}{\partial z} = f_1 + 2f_2z + 3f_3z^2 + \dots + (N-1)f_{N-1}z^{N-2} \quad (10)$$

108 To satisfy the boundary condition automatically, we expanded the velocity potential in a series of z^μ .

$$\begin{aligned} \phi(x, y, t) &= \sum_{\alpha=0}^{N-1} z^\alpha f_\alpha(x, y, t) \\ &= f_0 + z^\mu f_1 + z^{2\mu} f_2 + z^{3\mu} f_3 + \dots + z^{\mu(N-1)} f_{N-1} \end{aligned} \quad (11)$$

109 When $\mu > 1$, the vertical velocity can be given as Eq. (12), which satisfies the flat bottom boundary
 110 condition, $w = 0$ at $z = 0$. Also, when we analyze solitary waves, the vertical profile of velocity potential
 111 should include a low-order power function of z , such as less than cubic or fourth order. Therefore, it can be
 112 suggested that μ should be more than 1 and less than 3 or 4.

$$w = \frac{\partial \phi}{\partial z} = \mu f_1 z^{\mu-1} + 2\mu f_2 z^{2\mu-1} + 3\mu f_3 z^{3\mu-1} + \dots + \mu(N-1)f_{N-1}z^{\mu(N-1)-1} \quad (12)$$

113 Therefore, the system of new equations, which we call the modified FDS equation, is obtained by
 114 substituting Eq. (11) into Eqs. (6) and (7) as follows. It should be noted that the FDS equation can be given
 115 when $\mu = 1$.

$$\begin{aligned} \eta^{\mu\alpha} \frac{\partial \eta}{\partial t} + \frac{1}{\mu(\alpha + \beta) + 1} \nabla(\eta^{\mu(\alpha+\beta)+1} - b^{\mu(\alpha+\beta)+1}) \nabla f_\beta \\ - \frac{\mu^2 \alpha \beta}{\mu(\alpha + \beta) - 1} (\eta^{\mu(\alpha+\beta)-1} - b^{\mu(\alpha+\beta)-1}) f_\beta = 0 \end{aligned} \quad (13)$$

$$\eta^{\mu\beta} \frac{\partial f_\beta}{\partial t} + \frac{1}{2} \eta^{\mu(\beta+\gamma)} \nabla f_\beta \nabla f_\gamma + \frac{1}{2} \beta \gamma \eta^{\mu(\beta+\gamma)-2} f_\beta f_\gamma + g\eta + \frac{P}{\rho} = 0 \quad (14)$$

116 To solve Eqs. (13) and (14), an implicit technique was used in the numerical computational scheme by
 117 following Nakayama and Kakinuma (2010).

118

119 **2.3. Validation of modified FDS equation by solitary waves**

120 To verify the application of the modified FDS equation, we applied the model into solitary waves.
 121 Grimshaw (1971) and Fenton (1972) developed theoretical solutions, which reproduce surface solitary waves
 122 accurately. When a wave height and water depth are given as a_H and h , the wave speed is

$$c_K = \sqrt{gh} \left(1 + \frac{1}{2} \epsilon \right) \quad (15)$$

$$c_F = \sqrt{gh} \left(1 + \epsilon - \frac{1}{20} \epsilon^2 - \frac{3}{70} \epsilon^3 \right) \quad (16)$$

$$\epsilon = \frac{a_H}{h} \quad (17)$$

123 where c_K is the wave speed by the KdV theory, and c_F is the wave speed by Fenton (1972).

124 We made comparisons of not only wave speed but also an effective wavelength between the theory and
 125 model. The effective wavelength is defined as follows (Koop and Butler, 1981).

$$\frac{\lambda_K}{h} = 2 \sqrt{\frac{4h}{3a_H}} \quad (18)$$

$$\lambda_F = \int_{-\infty}^{\infty} \eta_F dx \quad (19)$$

$$\frac{\eta_F}{h} = 1 + \epsilon s^2 - \frac{3}{4} \epsilon^2 s^2 \tau^2 + \epsilon^3 \left(\frac{5}{8} s^2 \tau^2 - \frac{101}{80} s^4 \tau^2 \right) \quad (20)$$

$$s = \operatorname{sech} \left(\frac{\kappa x}{h} \right) \quad (21)$$

$$\tau = \tanh \left(\frac{\kappa x}{h} \right) \quad (22)$$

$$\kappa = \left(\frac{3}{4} \epsilon \right)^{1/2} \left(1 - \frac{5}{8} \epsilon + \frac{71}{128} \epsilon^2 \right) \quad (23)$$

126 where λ_K is the effective wavelength by the KdV theory, λ_F is the effective wavelength by Fenton (1972), and
 127 η_F is the surface displacement by Fenton (1972).

128

129 **2.4. Laboratory experiment**

130 The length, width and water depth of an open channel were set to 25.0 m, 1.0 m and 0.5 m, respectively
 131 (Fig. 2). A bump with a height of 0.25 m was introduced in the middle of the tank. The length of the bump at
 132 the bottom was 3.0 m, and the slope gradient was 1:5 on both sides with a width of 1.0 m. The bump consisted

133 of fine sand with a mean diameter of 0.21 mm. Surface waves were generated using a wave paddle at the left
134 end of Fig. 2, and a wave absorber was located at the right end of Fig. 2. We measured wave height at wave
135 gauge stations using wave gauges, and velocity was measured over a bump ($x=8.5$ m and $z=0.26$ m) using an
136 Acoustic Doppler Velocimeter (ADV). Wave height was measured at wg1 ($x=0$ m), wg2 ($x=0.55$ m), wg3
137 ($x=8.5$ m), wg4 ($x=7.0$ m), and wg5 ($x=9.0$ m), in which the wave height at wg1 was used as a boundary
138 condition for wave generation.

139 We carried out three laboratory experiments: Case 1 and Case 2 for waves without current, and Case 3 for
140 waves with current. It should be noted that each case has different measurements of wave height and velocities
141 (Table I). Wave height and wave period conditions were 0.10 m and 1.0 s (Case 1) and 0.14 m and 1.5 s (Case
142 2) for surface wave case without current. In Case 1, wave height was measured at wg2 and wg3, and horizontal
143 and vertical velocities were measured at wg3. In Case 2, wave height was measured at wg3 and wg4, and
144 horizontal velocity was measured at wg3. In addition to the above mentioned two cases, we carried out
145 laboratory experiments under the condition of 0.10 m and 1.0 s by giving a 0.02 m/s opposing current, which
146 is referred to as Case 3. In Case 3, wave height was measured at wg2, wg3 and wg5, and horizontal velocity
147 was measured at wg3.

148

149 3. Results

150 3.1. *Solitary waves*

151 For the validation of the modified FDS equation using the theoretical solutions of surface solitary waves,
152 the total number of horizontal meshes was given as 1000 with a mesh size of 0.04 m across the whole
153 computational domain, and a time step of 0.00005 s was used with a total time step of 20 s. CFL was estimated
154 to be about 360, which is small enough to carry out computations accurately. Regarding the value of N that
155 determines the computational accuracy, N was given as 3 by following Nakayama and Kakinuma (2010) and
156 Nakayama et al. (2019a) which shows that $N=3$ can provide accurate enough computational results even under
157 deep-water waves. Regarding the value of μ , we gave $\mu = 2.5$. Please see “4. Discussion” for more details,
158 including discussion of the effect of N on the computational accuracy. Seven different wave heights normalized
159 by total water depth were given from about 0.09 to 0.6 with an interval of 0.085. The wave speed of solitary

160 waves agrees well with the theory in Fenton (1972) and becomes slightly smaller than in the KdV theory with
 161 the increase in wave height (Solid lines in Fig. 3a). On the other hand, the larger the wave height, the larger
 162 the effective wavelength compared to the KdV theory (Solid lines in Fig. 3b). The modified FDS equation
 163 successfully reproduced not only the relationship between the normalized wave height and wave speed but
 164 also the relationship between the normalized wave height and effective wavelength (Fig. 3). Therefore, the
 165 validity of the modified FDS equation was approved through comparisons with theoretical solutions.

166

167 **3.2. Waves without current**

168 The total number of horizontal meshes was given as 500 with a mesh size of 0.05 m across the whole
 169 computational domain, and a time step of 0.0005 s was used with a total time step of 50 s. A timeseries of
 170 wave heights at the left end was given using a timeseries of measured wave height at wg1. To give similar
 171 conditions as the laboratory experiments, a wave absorber was given at the right end with a length of 10 m in
 172 order to prevent progress waves from reflecting. N was given as 3 the same as solitary wave case. Since fine
 173 sediment with a diameter of 0.21 mm was used at the bottom of the tank, bottom friction loss of momentum
 174 was considered in the numerical computations (Eq.(24)). In this study, $f=0.052$ was decided on by changing
 175 the value of f by trial and error.

$$momentum\ loss = -\rho \frac{f_c}{h} |u_b| u_b \quad (24)$$

176 where, u_b is the velocity at the bottom, and f_c is the coefficient of energy loss at the bottom.

177 In the computation, 5 different values of μ was applied in the modified FDS equation, 1.0, 1.5, 2.0, 2.5
 178 and 3.0. But, the larger the μ given, the more unstable the numerical computation becomes. In Case 2, the
 179 condition of $\mu = 3.0$ was very unstable and we could not obtain the results. Therefore, there are 4 different
 180 conditions, $\mu = 1.0, 1.5, 2.0$ and 2.5 , in Case 2.

181 In Case 1, the wave height at wg2 agrees very well for all the computational results because wg2 is located
 182 close to wave generator wg1 (Figs. 4 and 5). In contrast to wg2, the FDS equation slightly overestimated wave
 183 height at wg3 when $\mu = 1.0$, and the modified FDS equation underestimated wave height less than $\mu = 1.0$
 184 when $\mu = 3.0$. However, it can be said that wave heights from all the computational cases agree well with
 185 laboratory experiments. On the other hand, there were large differences between numerical computations and

laboratory experiments for velocity at wg3. When $\mu = 1.0$ (corresponding to the FDS equation), horizontal velocity was overestimated by a factor of two compared to the laboratory experiments (Fig. 4). However, when $\mu > 1.0$ (corresponding to the modified FDS equation), horizontal velocity was estimated to be close to the laboratory experiments, with best agreement when $\mu = 2.5$ (Fig. 5). Interestingly, opposite to horizontal velocity, vertical velocity was underestimated, with the velocity 1/3 of the laboratory experiments when $\mu = 1.0$ (the FDS equation). When $\mu > 1.0$ (the modified FDS equation), vertical velocity estimates were close to the laboratory experiments with the best estimation obtained when $\mu = 2.5$ and vertical velocity overestimated when $\mu = 3.0$.

In Case 2, where the wave height is larger and wave period longer than Case 1, wave height was slightly overestimated at wg3 and wg4 when $\mu = 1.0$ (the FDS equation), which is like Case 1 (Fig. 6). When $\mu > 1.0$ (the modified FDS equation), wave height was estimated very well and the best estimation was found when $\mu = 2.5$ (Fig. 7). Similar to Case 1, horizontal velocity was overestimated at wg3 largely when $\mu = 1.0$, but horizontal velocity was estimated close to the laboratory experiments when $\mu > 1.0$. The best agreements were found when $\mu = 2.0$ or $\mu = 2.5$. Therefore, there may be the possibility that the FDS equation overestimates horizontal velocity and underestimates vertical velocity when $\mu = 1.0$. On the other hand, when $\mu > 1.0$ (the modified FDS equation), horizontal and vertical velocities were estimated very well, with the best agreement when $\mu = 2.5$. This may thus suggest that the application of the modified FDS equation is recommended.

203

204 **3.3. Waves with current**

In the analysis for wave-current conditions, it was considered appropriate to apply $\mu = 2.5$ based on the results obtained under the wave condition without current shown in the previous section. Therefore, we applied $\mu = 2.5$ into Case 3 under conditions of wave heights of 0.10 m and the wave periods of 1.0 s against a current of 0.02 m/s. For all the wave gauge stations, wave heights at wg2, wg3 and wg5 were found to agree with the laboratory experiments very well (Fig. 8). Also, horizontal velocity was estimated through the numerical computation, although it gave slightly larger velocities. In comparison with Case 1 (wave without current but the wave height and wave period conditions are the same as Case 3), wave heights under Case 3 at wg3 over

212 a bump were found to be smaller due to the current effect. It is apparent that the amplitude of horizontal
 213 velocity was smaller than Case 1 while the mean horizontal velocity was confirmed to be 0.02 m/s smaller
 214 than Case 1, which is the effect of the constant 0.02 m/s opposing current.

215

216 4. Discussion

217 For evaluating the modified FDS equation, the linearized modified FDS equation were obtained to
 218 investigate the dispersion relationship from the viewpoint of changes in μ and N . Eqs. (25) and (26) can be
 219 obtained by linearizing Eqs. (13) and (14).

$$H^{\mu\alpha} \frac{\partial \eta}{\partial t} + \frac{H^{\mu(\alpha+\beta)+1}}{\mu(\alpha+\beta)+1} \frac{\partial^2 f_\beta}{\partial x^2} - \frac{\mu^2 \alpha \beta}{\mu(\alpha+\beta)-1} H^{\mu(\alpha+\beta)-1} f_\beta = 0 \quad (25)$$

$$H^{\mu\beta} \frac{\partial f_\beta}{\partial t} + g\eta = 0 \quad (26)$$

220 Eq. (27) is obtained from Eqs. (25) and (26).

$$-\frac{\partial^2 f_\beta}{\partial t^2} + \frac{gH}{\mu(\alpha+\beta)+1} \frac{\partial^2 f_\beta}{\partial x^2} - \frac{\mu^2 \alpha \beta}{\mu(\alpha+\beta)-1} \frac{g}{H} f_\beta = 0 \quad (27)$$

221 The following relationships are obtained by applying small amplitude surface waves.

$$f_\beta = G_\beta \exp[\sigma t - kx] \quad (28)$$

222 Substituting Eq. (28) into Eq. (27) provides Eq. (29).

$$\sigma^2 - \frac{gH}{\mu(\alpha+\beta)+1} k^2 - \frac{\mu^2 \alpha \beta}{\mu(\alpha+\beta)-1} \frac{g}{H} = 0 \quad (29)$$

223 where, k is the wave number, and σ is the frequency.

224 The theoretical solution of the dispersion relationship for surface waves, Eq. (30), was compared with the
 225 linearized modified FDS equation, Eq. (29) (Fig. 9). We applied 9 different μ ($=1.0, 1.5, 2.0, 2.5, 3.0, 3.5, 4.0,$
 226 $4.5, 5.0$) and 5 different N ($=1, 2, 3, 4, 5$).

$$c^2 \frac{1}{c_0^2} = \left(\frac{\sigma}{k}\right)^2 \frac{1}{gh} = \frac{1}{kh} \tanh(kh) \quad (30)$$

227 When $\mu = 1.0$ (the FDS equation), solitary-like wave conditions were found to be satisfied by giving $N =$
 228 3 ($0 < kh < 1$). However, N needs to be larger than 3 to satisfy deep water wave conditions ($\mu = 1.0$). For μ
 229 $= 1.5$ and $\mu = 2.0$, the lines of $N = 2$ become closer to the theoretical solution compared to $\mu = 1.0$ though N also
 230 needs to be larger than 3 to satisfy deep water wave conditions. Interestingly, when $\mu = 2.5$, the lines of $N = 3$
 231 agree with the theoretical solution very well. The lines when N is larger than 2 for $\mu = 2.5$ agree with the
 232 theoretical solutions for not only shallow water conditions but also deep-water conditions. Comparison with
 233 the laboratory experiments demonstrated that the computational results agree very well when $\mu = 2.5$, which
 234 may suggest that the best recommended value of μ may be 2.5.

235 On the other hand, when $\mu = 3.0$, results appeared similar to $\mu = 2.5$. However, there are slight gaps when
 236 kh was less than 1.0. When μ was more than 3.0, it was more apparent that Eq. (29) cannot agree with the
 237 theoretical solution when kh is less than 1.0, though deep-water conditions are satisfied even when $N = 2$. It is
 238 thus revealed that the larger the value of μ , the more the deep-water conditions are satisfied (though shallow
 239 water and solitary-like wave conditions cannot be satisfied). Therefore, from the viewpoint of the dispersion
 240 relationship and comparisons with laboratory experiments, $\mu = 2.5$ is the recommended value for surface wave
 241 computations.

242 For Case 3 where currents were given against progressive waves, the wave height at wg2 was smaller than
 243 in Case 1, which had the same conditions for wave height and wave period without currents. Also, the wave
 244 height at wg3 in Case 3 was smaller than Case 1. Dean (1965) and Dalrymple (1974) showed the significant
 245 effect of currents on the profile of a surface wave. Since currents against progressive waves were given in Case
 246 3, wave speed in Case 3 is smaller than Case 1, which suggests that wave height becomes smaller resulting in
 247 an increase in effective wavelength. Such a larger effective wavelength wave may be categorized as a “bulbous”
 248 wave.

249 In previous studies, when a linear shear or constant opposing current are applied to progressive waves,
 250 Teles and Peregrine (1988) suggested that a “bulbous” wave that having a thick or a larger width ridge may

251 occur. Therefore, as a fundamental analysis, we made an attempt to investigate whether a bulbous wave exists
252 or not using Case 3 where currents are given against progressive waves (Fig. 10). The wavelength of Case 3
253 was about 1.5 m base on the laboratory experiments and numerical computations, which shows that the wave
254 type is a deep-water wave, $kh = 2.1$ ($kh / (2\pi) = 0.33 < 1/2$) (Mei, 1989). Therefore, we made comparisons
255 with the theoretical solution of a Trochoidal wave which is one of the finite amplitude deep-water waves. It
256 appears that the profile of surface waves computed with the modified FDS equation has a larger width ridge
257 in comparison to the Trochoidal wave (red circle in Fig. 10). However, the apparent bulbous wave only
258 occurred in the region behind the bump, and we could not find a clear bulbous-type wave between the wave
259 generator and the bump. We also investigated only one wave-current interaction case. Therefore, it should be
260 noted that further investigation is necessary to confirm the occurrence of a bulbous wave.

261

262 5. Conclusion

263 The theoretical solution of the dispersion relationship for surface waves was compared with the linearized
264 modified FDS equation, which demonstrated the high applicability of the FDS equation when $\mu = 2.5$. Also,
265 N is recommended to be more than 3 when $\mu = 2.5$ in order to reproduce deep-water waves. By giving $\mu = 2.5$
266 and $N = 3$, the modified FDS equation was verified from comparisons with the theoretical solution by Fenton
267 (1972). The modified FDS equation was successfully applied in two laboratory experiments without currents
268 and one laboratory experiment with currents. It was found from the comparisons between with and without
269 current cases that wave height and effective wavelength become smaller and larger, respectively, when a
270 constant opposing current are applied to progressive waves.

271

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276 **References**

- 277 Agnon, Y., Madsen, P. A., Schaffer, H. A. 1999. "A new approach to high-order Boussinesq models." *J*
278 *Fluid Mech* 399: 319–333.
- 279 Dalrymple, R.A. 1974. "A finite amplitude wave on a linear shear current." *J Geophys Res.* 79: 4498-4504.
- 280 Dean, R.G. 1965. "Stream function representation of nonlinear ocean waves." *J Geophys Res.* 70 (18): 4561-
281 4572.
- 282 Grimshaw R. 1971. "The solitary wave in water of variable depth. Part 2." *J Fluid Mech* 46: 611-622.
- 283 Fenton J. 1972. "A ninth-order solution for the solitary wave." *J Fluid Mech* 53: 257-271.
- 284 Horn, D. A., Redekopp, L.G., Imberger, J., Ivey, G.N. 2000. "Internal wave evolution in a space-time
285 varying field." *J Fluid Mech* 424: 279-301.
- 286 Horn, D.A., Imberger, J., Ivey, G.N. 2001. "The degeneration of large-scale interfacial gravity waves in
287 lakes." *J Fluid Mech* 434: 181–207.
- 288 Horn, D.A., Imberger, J., Ivey, G.N., Redekopp, L.G. 2002. "A weakly nonlinear model of long internal
289 waves in closed basins." *J Fluid Mech* 467: 269-287.
- 290 Isobe, M. 1995. "Time-dependent mild-slope equations for random waves." *Proc 24th Int Conf Coast Eng*
291 *ASCE*: 285-299.
- 292 Koop, C.G., Butler, G. 1981. "An investigation of internal solitary waves in a two-fluid system." *J Fluid*
293 *Mech* 112: 225–254.
- 294 Madsen, O.S., Mei, C.C. 1969. "The transformation of a solitary wave over an uneven bottom." *J Fluid*
295 *Mech* 39(4): 781-791.
- 296 Madsen, P.A., Murray, R., Sorensen, O.R. 1991. "A new form of Boussinesq equations with improved linear
297 dispersion characteristics." *Coast Eng* 15: 371-388.
- 298 Madsen, P.A., Sorensen, O.R. 1992. "A new form of Boussinesq equations with improved linear dispersion
299 characteristics. Part 2. A slowly varying bathymetry." *Coast Eng* 18: 183-204.
- 300 Maxworthy, T. 1980. "On the formation of nonlinear internal waves from the gravitational collapse of mixed
301 regions in two and three dimensions." *J Fluid Mech* 96: 47-64.
- 302 Mei, C.C. 1989. "The applied dynamics of ocean surface waves. Advanced Series on Ocean Engineering."
303 *World Scientific* 1: 768.

304 Mei, C.C., Le, Mehaute, B. 1966. "Note on the Equations of Long Waves Over an Uneven Bottom." *J Geophys*
305 *Res* 71(2): 393-400.

306 Li, W., Yeh, H., Kodama, Y. 2011. "On the Mach reflection of a solitary wave: revisited." *J Fluid Mech* 672:
307 326-357.

308 Luke, J.C. 1967. "A variational principle for a fluid with a free surface." *J Fluid Mech* 27: 395-397.

309 Miles, J.W. 1997a. "Obliquely interacting solitary waves." *J Fluid Mech* 79: 157-169.

310 Miles, J.W. 1997b. "Resonantly interacting solitary waves." *J Fluid Mech* 79: 171-179.

311 Nakayama, K. 2006. "Comparisons of using CIP, compact and CIP-CSL2 schemes for internal solitary waves."
312 *Inter. J. Num. Method. Fluid.* 51: 197-219.

313 Nakayama, K., Kakinuma, T. 2010. "Internal waves in a two-layer system using fully nonlinear internal-
314 wave equations." *Int J Numer Meth Fluids* 62: 574-590.

315 Nakayama, K., Shintani, T., Kokubo, K., Maruya, Y., Kakinuma, T., Komai, K. and Okada, T. 2012.
316 "Residual current over a uniform slope due to breaking of internal waves in a two-layer system." *J.*
317 *Geophy. Res.* 117, C10002. doi: 10.1029/2012JC008155.

318 Nakayama, K., Kakinuma, T., Tsuji, H. 2019a. "Oblique reflection of large internal solitary wave." *Eur. J.*
319 *Mech. (B)*, Vol.74: 81-91.

320 Nakayama, K., Lamb, K. 2019c. "Breathers in a three-layer fluid." *J Fluid Mech* submitted.

321 Nakayama, K., Sato, T., Shimizu, K. and L. Boegman. 2019b. "Classification of internal solitary wave
322 breaking over a slope." *Physical Review Fluids*, Vol.4, 014801

323 Nwogu, O. 1993. "An alternative form of the Boussinesq equations for nearshore wave propagation." *J*
324 *Waterw Port Coast Ocean Eng.* 119(6): 618-638.

325 Perigrine, D.H. 1967. "Long waves on a beach." *J Fluid Mech* 27(4): 815-827.

326 Tanaka, M. 1993. "Mach reflection of a large-amplitude solitary wave." *J Fluid Mech* 248: 637-661.

327 Teles, da, Silva, A.F., Peregrine, D. 1988. "Steep steady surface waves on water of finite depth with constant
328 vorticity." *J Fluid Mech* 195: 281-302.

329 Thuy, Thi, Thu, Vu. 2013. "Aspects of inlet geometry and dy-namics." *The University of Queensland, Ph. D.*
330 *Thesis.*

331 Wei, G., Kirby, J.T. 1995. "A time-dependent numerical code for the extended Boussinesq equations." *J*
332 *Waterw Port Coast Ocean Eng* 121(5): 251-261.

333 Wei, G., Kirby, J.T., Grilli, S.T. and Subramanya, R. 1995. "A fully nonlinear Boussinesq model for surface
334 waves. Part 1. Highly nonlinear unsteady waves." *J Fluid Mech* 294: 71–92.

335 Wei, G., Kirby, J.T., Sinha, A. 1999. "Generation of waves in Boussinesq models using a source function
336 method." *Coast Eng* 36: 271-299.

337 Zaman, M.H., Hirayama, K., Hiraishi, T. 2000. "A Boussinesq Model to Study Long Period Waves in a
338 Harbor." *Rept of Port and Harbour Res. Inst.* 39(4): 25-50.

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Table 1 Conditions of laboratory experiment.

	wave height (m)	wave period (s)	current (m/s)	wg1	wg2	wg3	wg4	wg5	ADV horizontal velocity (m/s)	ADV vertical velocity (m/s)
Case 1	0.10	1.0	0	O	O	O	--	--	O	O
Case 2	0.14	1.5	0	O	--	O	O	--	O	--
Case 3	0.10	1.0	-0.02	O	O	O	--	O	O	--

O: measured. --: not measured.

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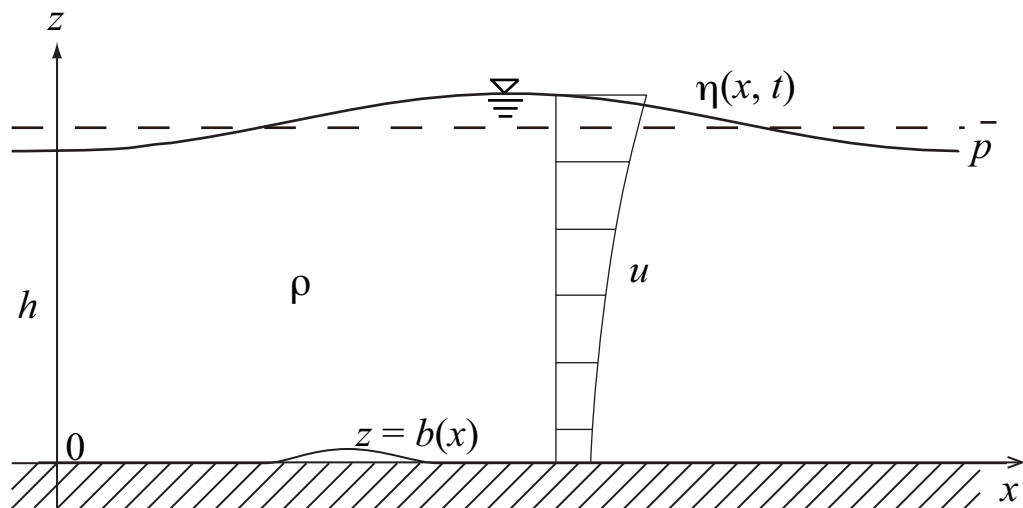


Fig. 1 Schematic diagram of surface waves for the FDS equation.

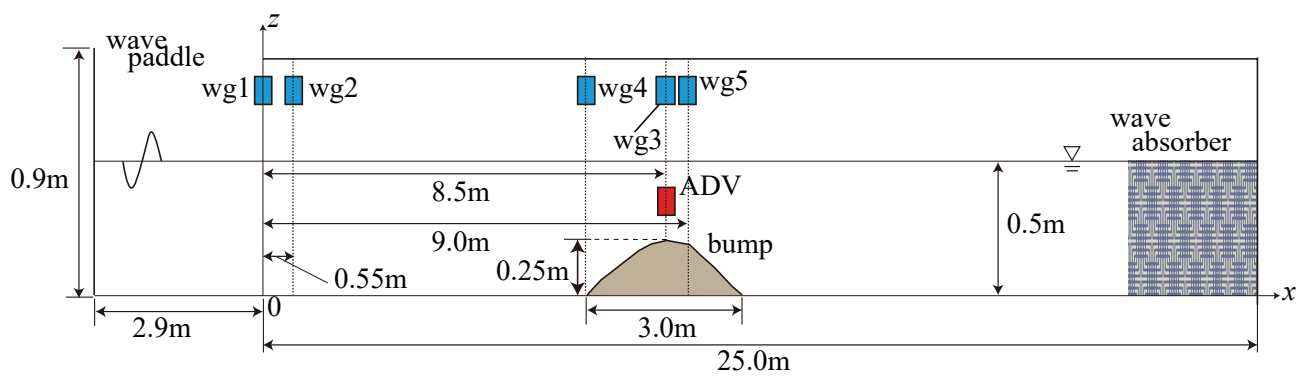


Fig. 2 Open channel tank for laboratory experiment and wave gauge stations and the location of ADV.

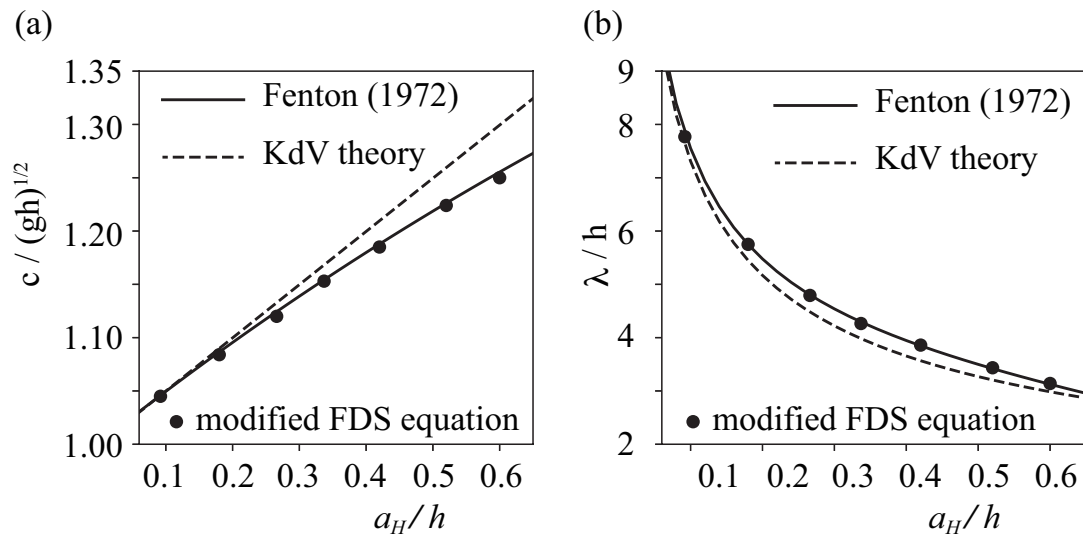


Fig. 3 Comparisons between theoretical solutions and the modified FDS model. (a) Normalized wave height and wave speed. (b) Normalized wave height and wavelength.

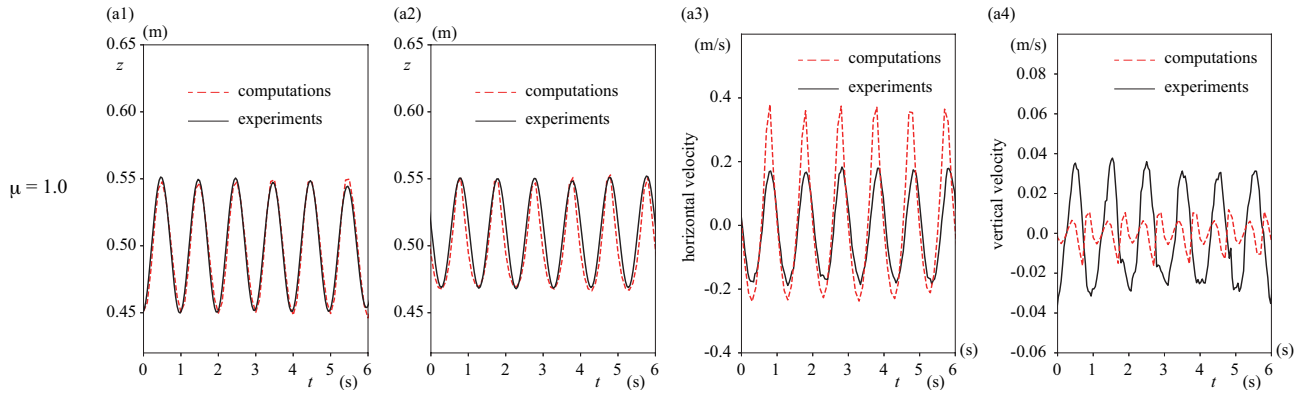


Fig. 4 Comparisons with laboratory experiment case 1 when $\mu = 1.0$. (a1) wave amplitude at wg2. (a2) wave amplitude at wg3. (a3) horizontal velocity at wg3. (a4) vertical velocity at wg3.

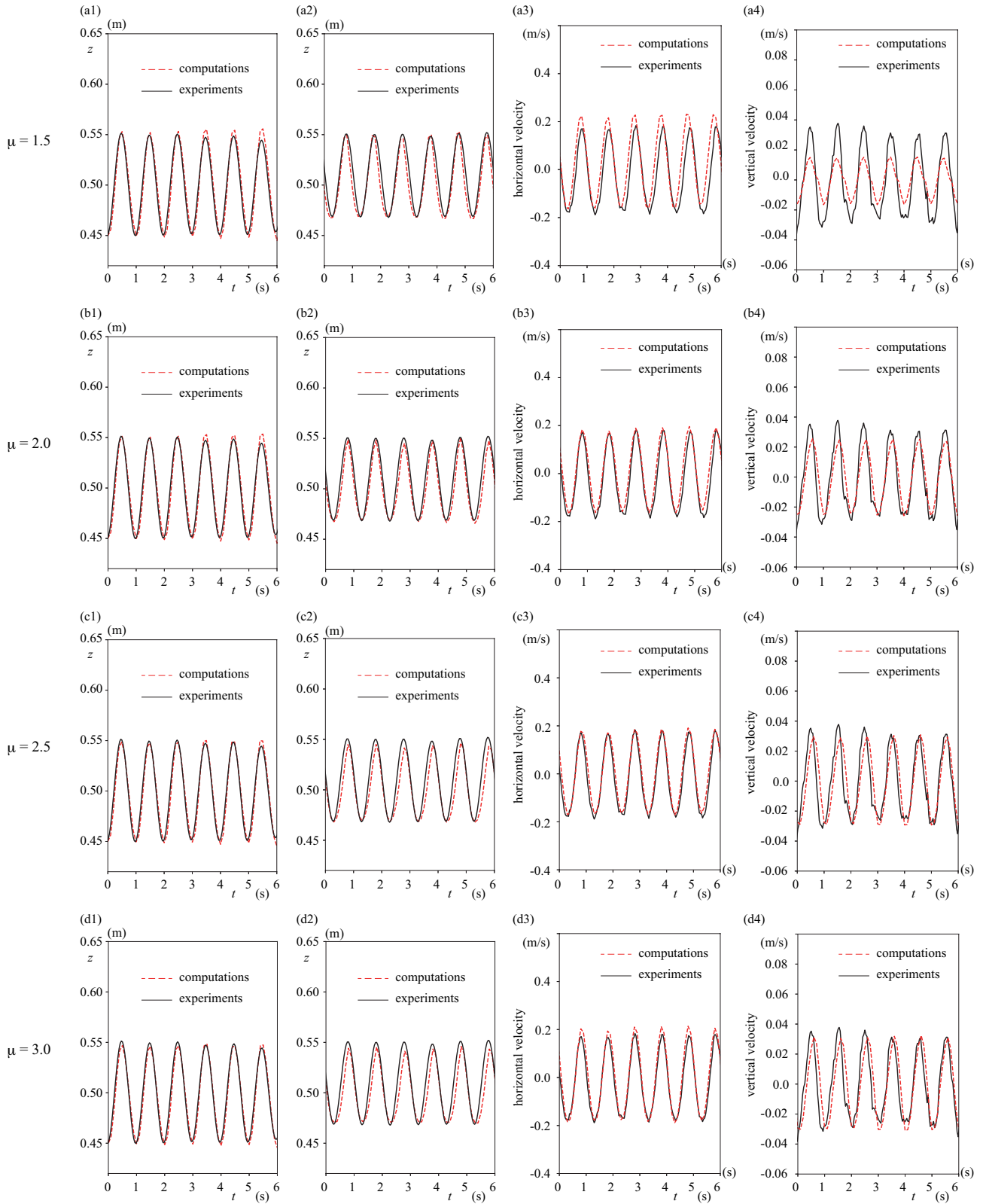


Fig. 5 Comparisons with laboratory experiment case 1 when (a) $\mu = 1.5$, (b) $\mu = 2.0$, (c) $\mu = 2.5$, and (d) $\mu = 3.0$. (a1-d1) wave amplitude at wg2. (a2-d2) wave amplitude at wg3. (a3-d3) horizontal velocity at wg3. (a4-d4) vertical velocity at wg3.

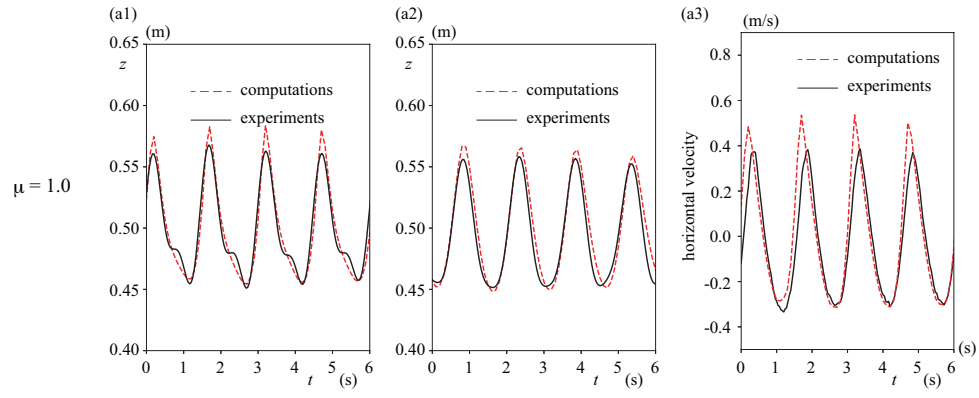


Fig. 6 Comparisons with laboratory experiment case 2 when $\mu = 1.0$. (a1) wave amplitude at wg3. (a2) wave amplitude at wg4. (a3) horizontal velocity at wg3.

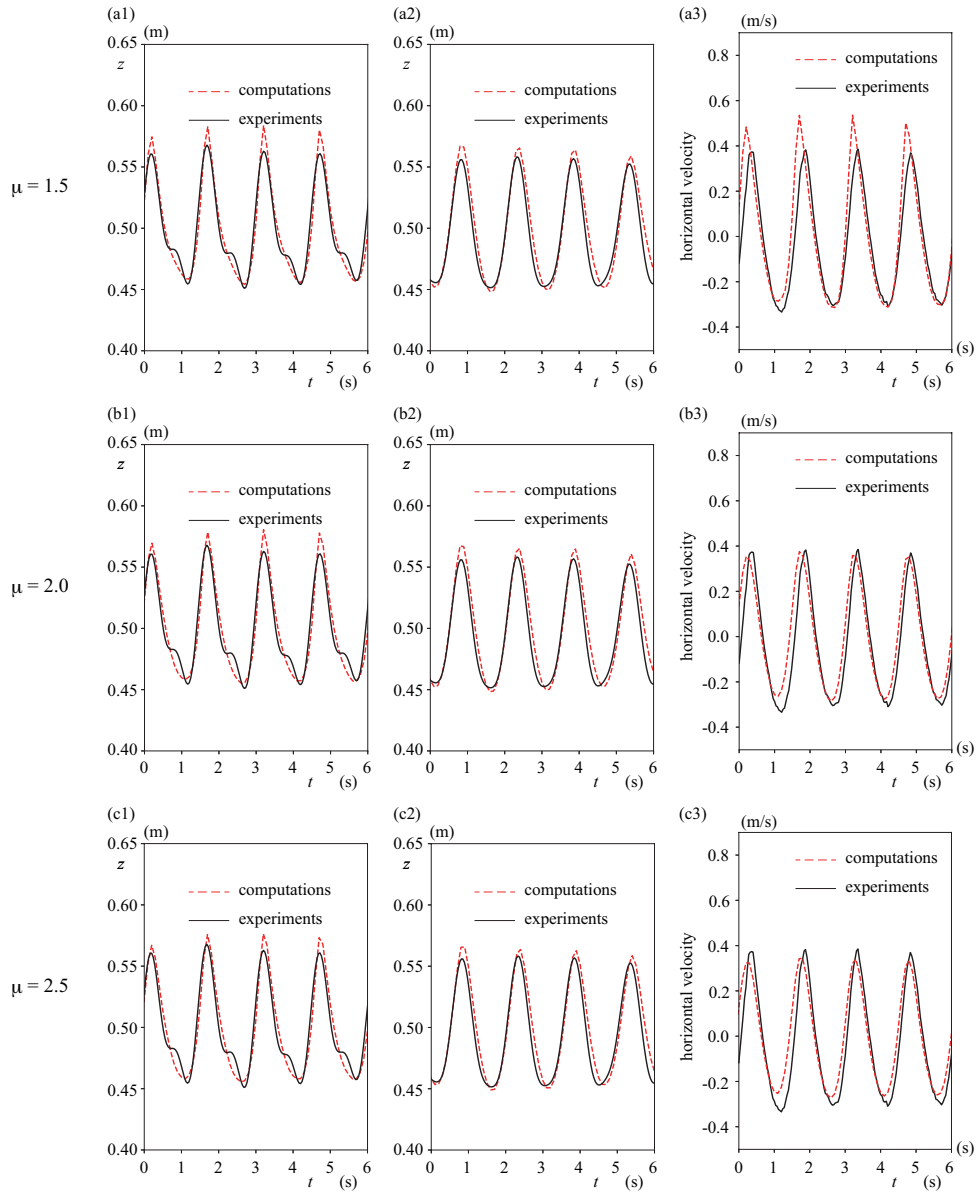


Fig. 7 Comparisons with laboratory experiment case 1 when (a) $\mu = 1.5$, (b) $\mu = 2.0$, and (c) $\mu = 2.5$. (a1-c1) wave amplitude at wg3. (a2-c2) wave amplitude at wg4. (a3-c3) horizontal velocity at wg3.

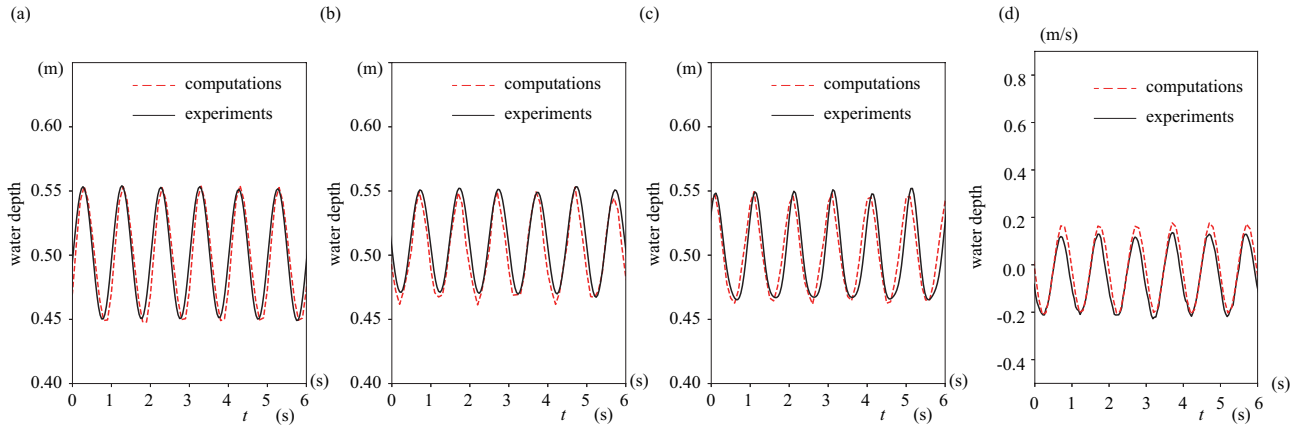


Fig. 8 Comparisons with laboratory experiment case 3 when $\mu = 2.5$. (a) wave amplitude at wg2. (b) wave amplitude at wg3. (c) wave amplitude at wg5. (d) horizontal velocity at wg3.

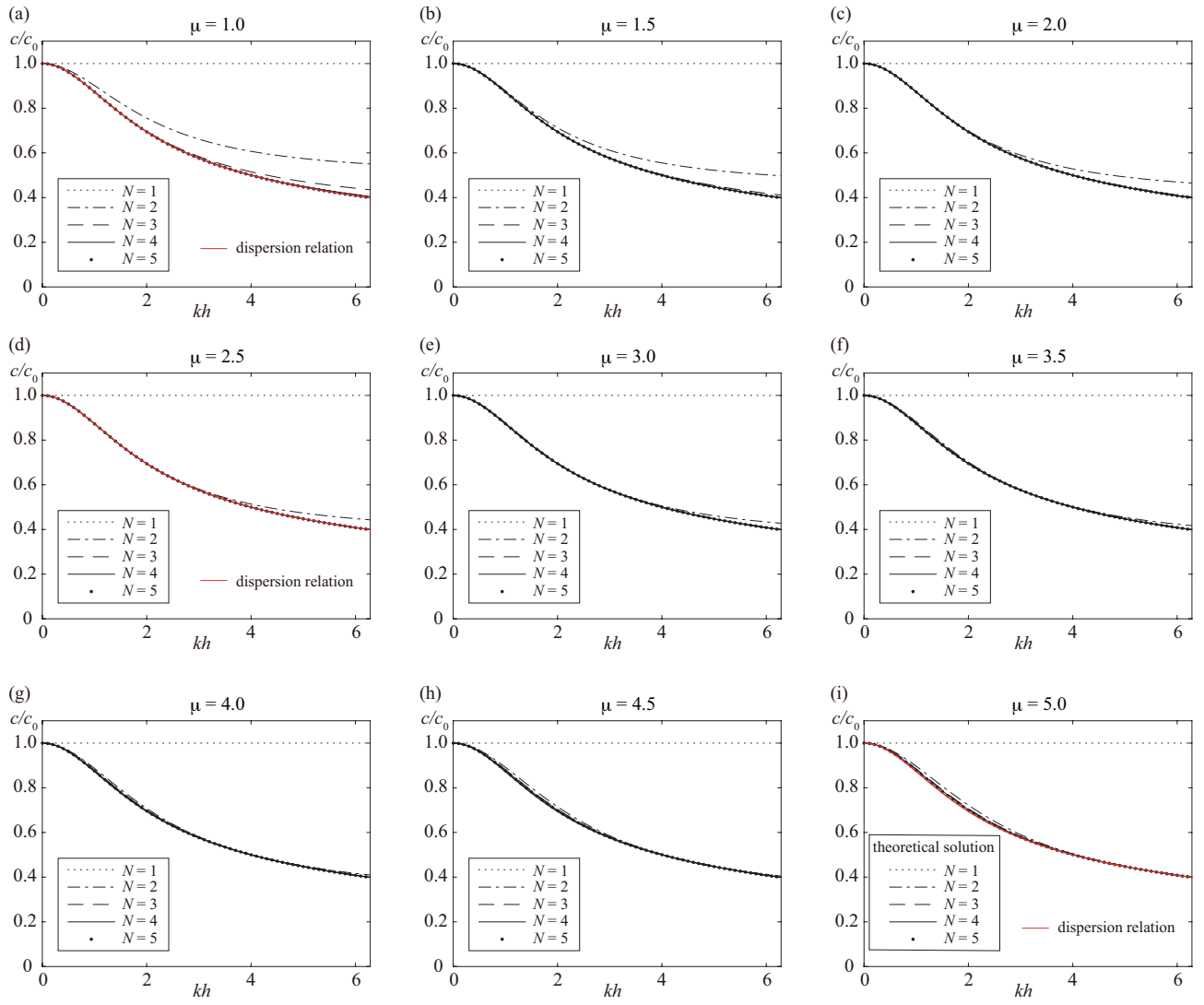


Fig. 9 Dispersion relationship. Red lines indicate the theoretical solution. (a) $\mu = 1.0$, (b) $\mu = 1.5$, (c) $\mu = 2.0$, (d) $\mu = 2.5$, (e) $\mu = 3.0$, (f) $\mu = 3.5$, (g) $\mu = 4.0$, (h) $\mu = 4.5$, and (i) $\mu = 5.0$.

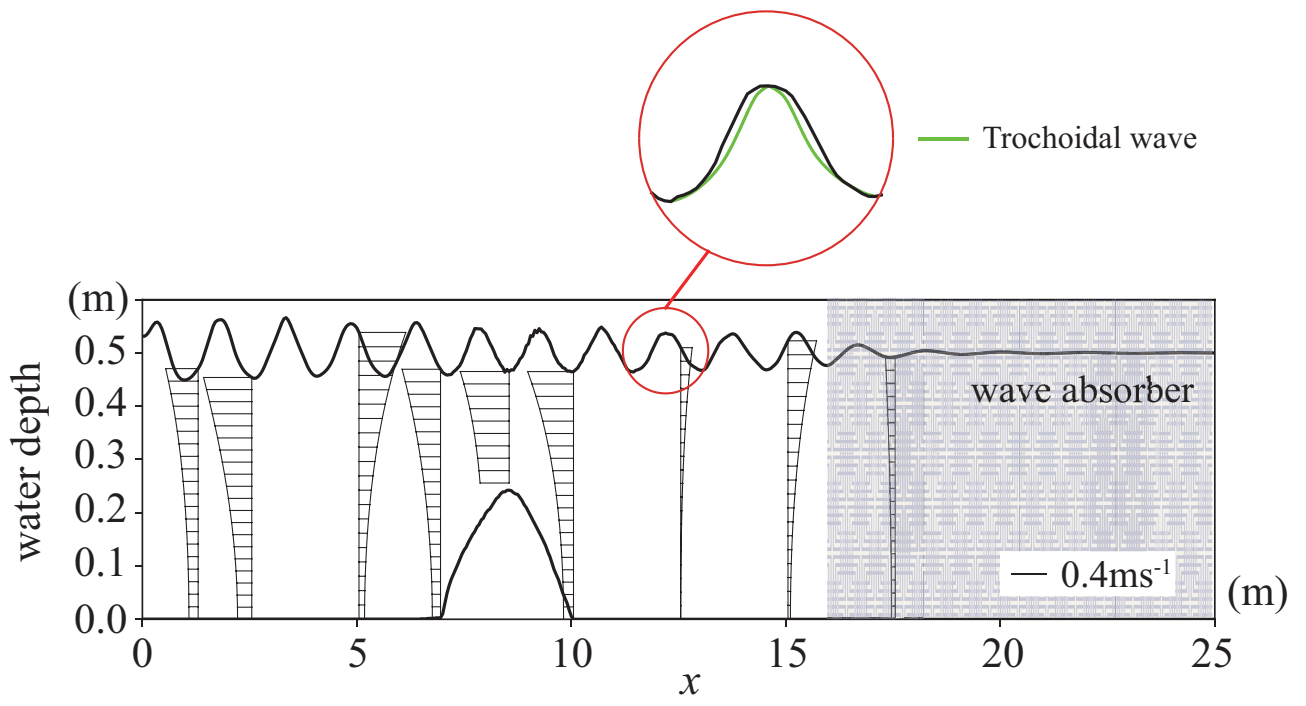


Fig. 10 Wave profile and vertical profile of horizontal velocity in case 3.